

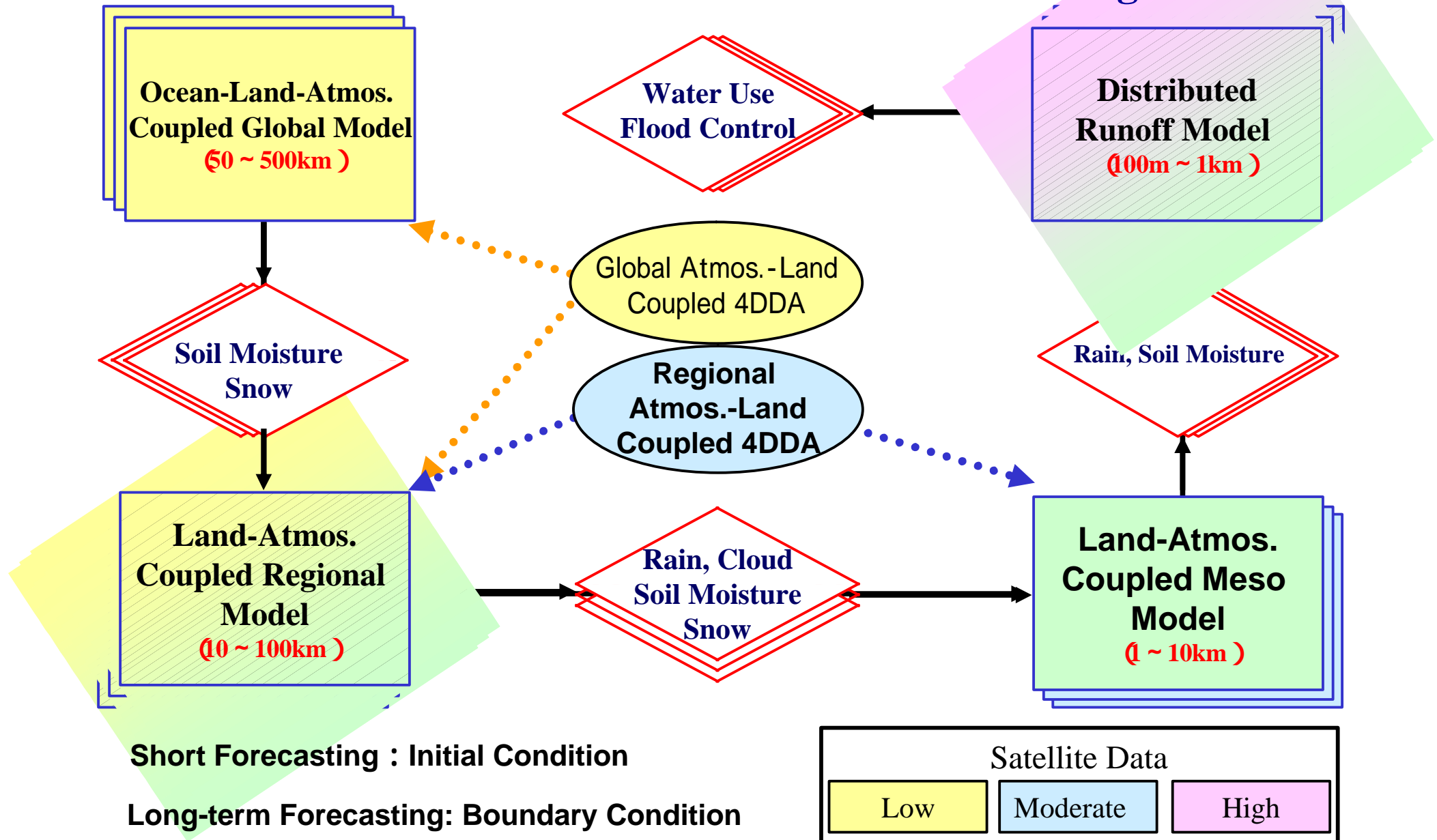
An auto-calibration system  
to assimilate AMSR data into a land surface model  
for estimating soil moisture and surface energy budget

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H. FUJII<sup>1</sup>, K. TAMAGAWA<sup>1</sup>, Y. MA<sup>3</sup>, H. ISHIKAWA<sup>4</sup>

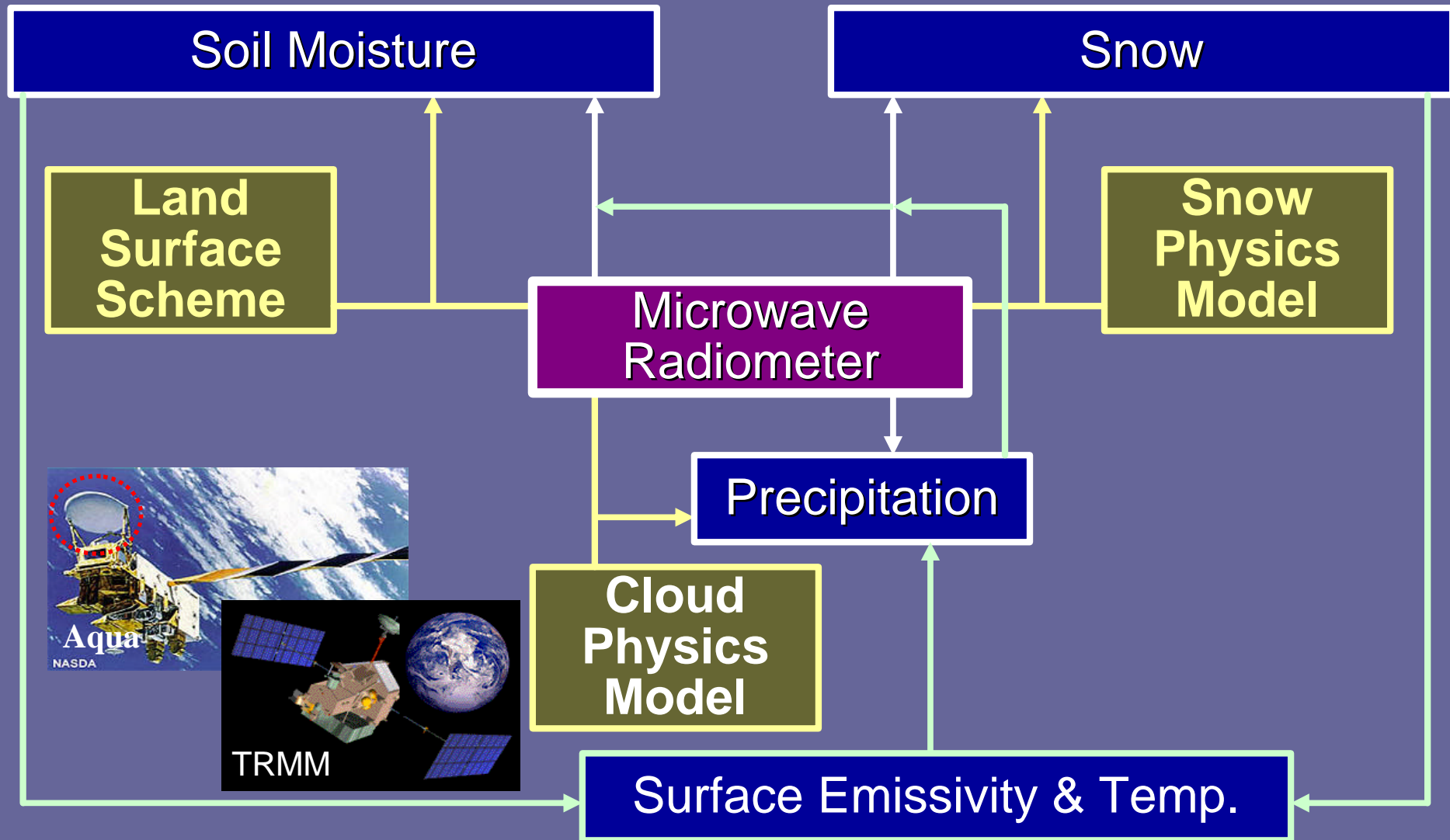
1. The University of Tokyo
2. CAREERI, Chinese Academy of Sciences
3. Inst. Tibetan Plateau, Chinese Academy of Sciences.
4. Kyoto University

Global      Regional-Meso      Basin

# Satellite Data Assimilation and Down-scaling



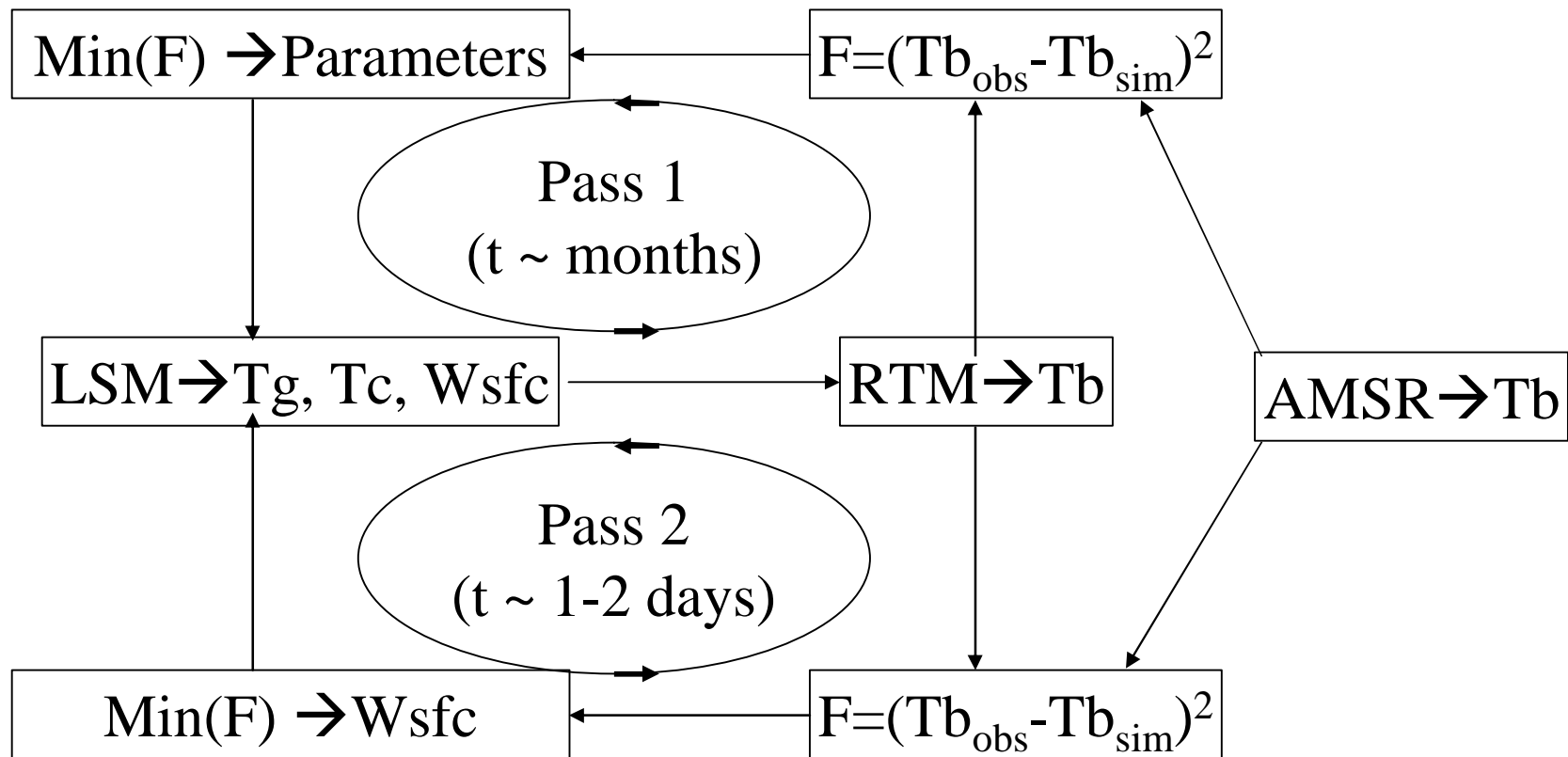
# Four Dimensional Data Assimilation



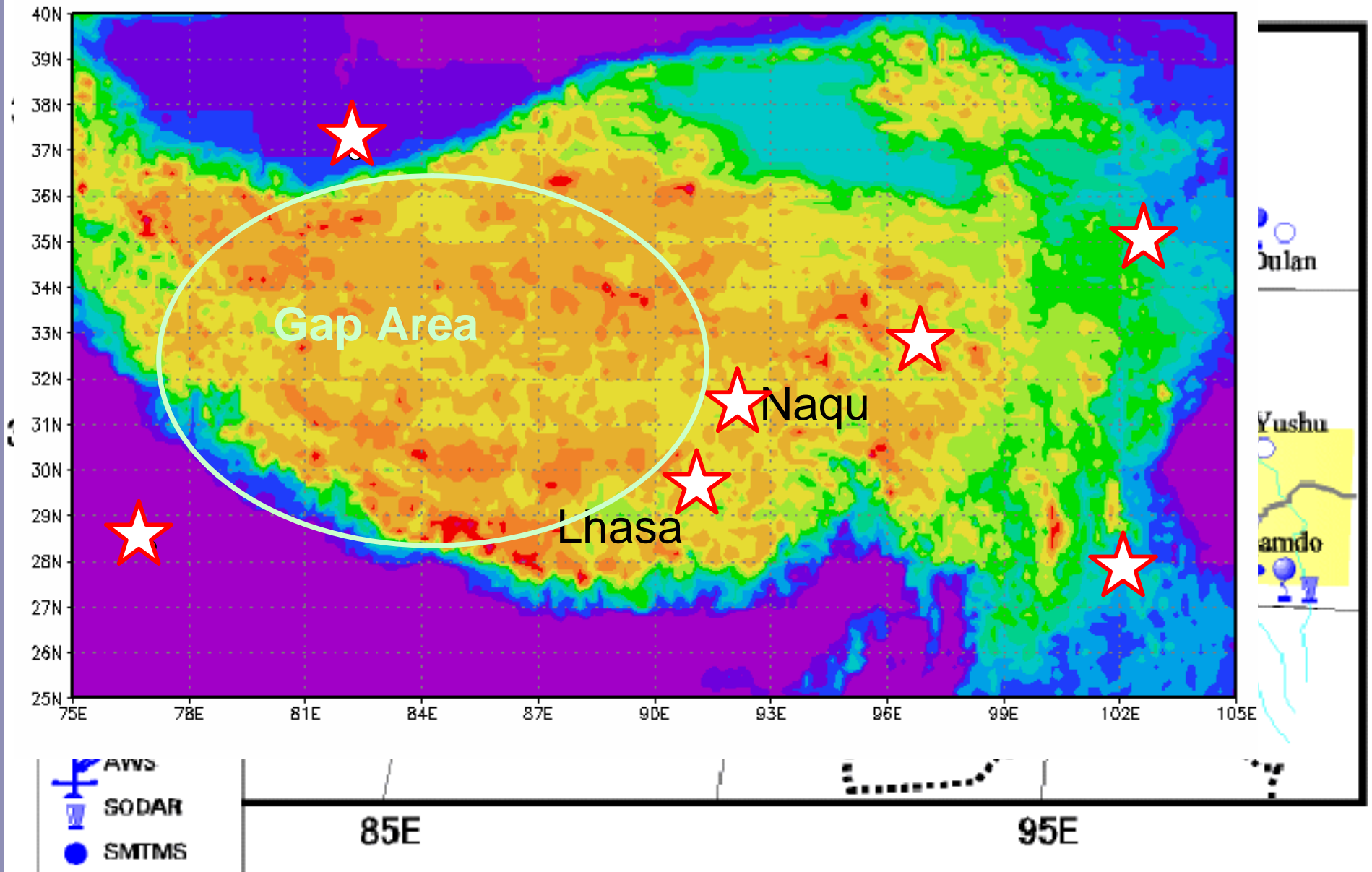
# Two-pass LDAS-UT

soil texture, porosity: LSM

surface roughness: RTM

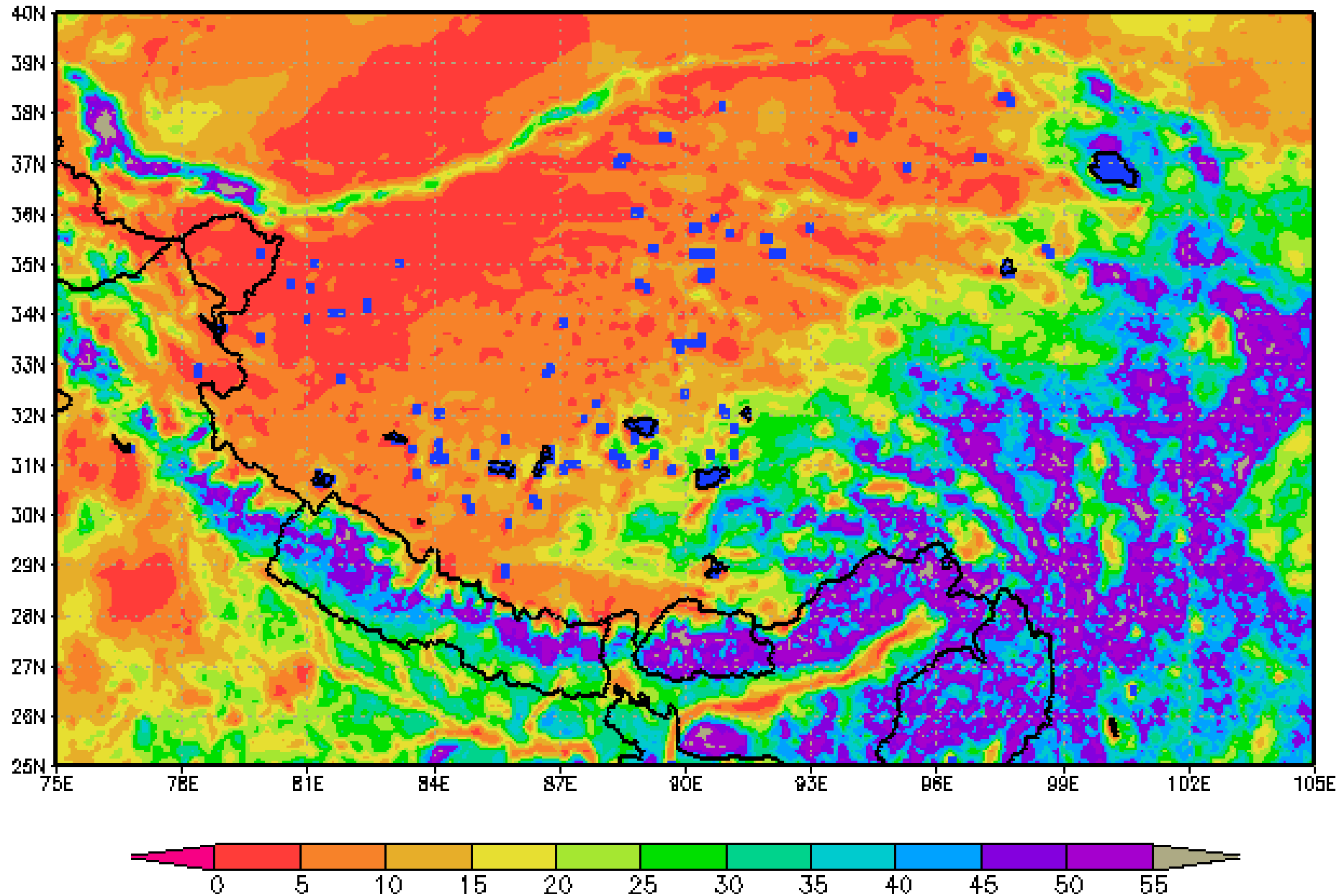


# Plateau Scale Experiment



# Seasonal Variation of the Soil Moisture in the Tibetan Plateau

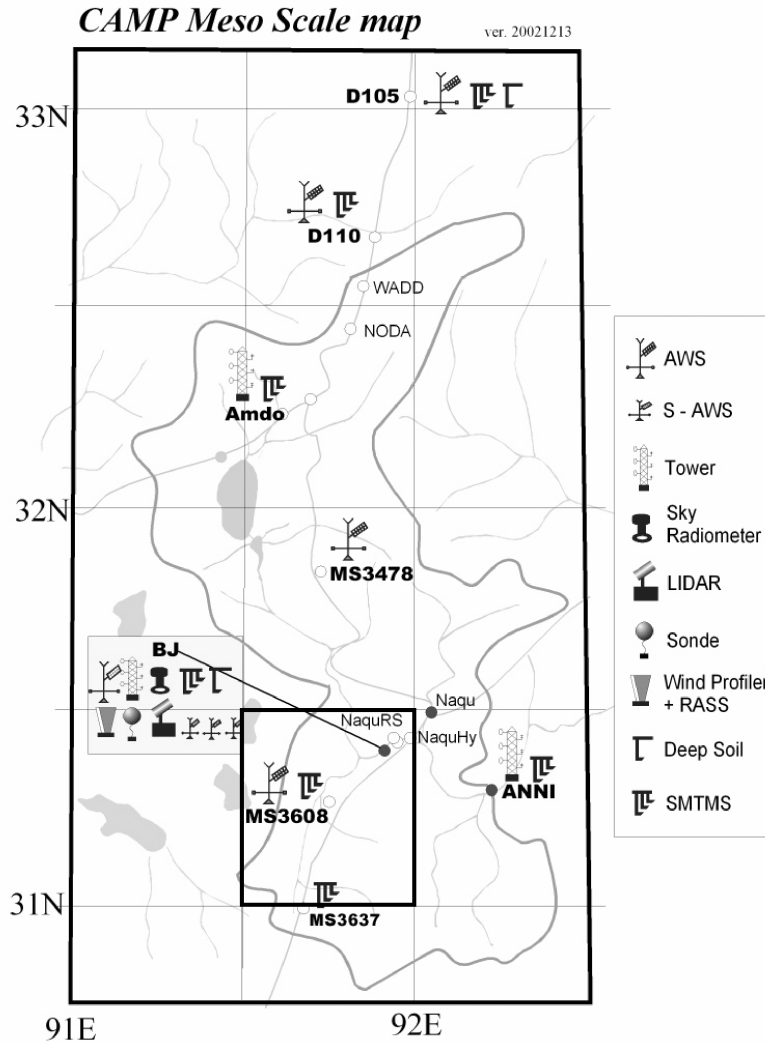
6G Mv(%) tibet\_D 2003SEP-last



## **Input data → Easy application in any region**

- LDAS-UT grid size: 0.5 degree
- Forcing
  - GPCP precipitation: 1 degree
  - ISCCP radiation: 2.5 degree
  - NCEP reanalysis: 1.5 degree
- Leaf area index: MODIS 0.25 degree 8-day product
- Microwave Tb: AMSR 0.5 degree 6.9 and 19.7 GHz

# First application: A case at CEOP Tibet site



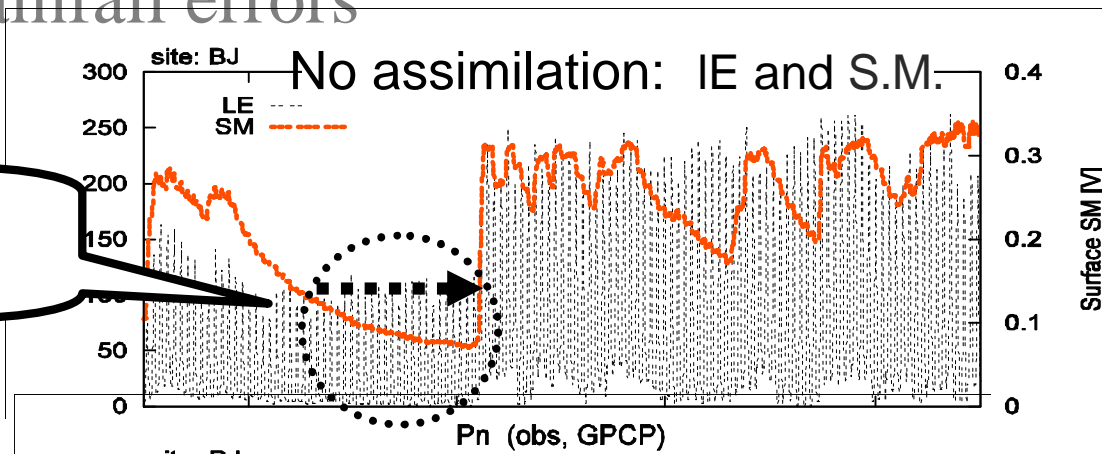
| Items               | Station (depth)                            |
|---------------------|--|
| Precipitation       | BJ   |
| Radiation           | BJ   |
| Surface temperature | BJ, MS3608                                 |
| Near-surface        | BJ, MS3608 (4cm)                           |
| soil moisture       | S-AWS1, S-AWS3 (0-5 cm)<br>SSMTMS (0-3 cm) |
| Turbulent fluxes    | BJ (3m, 20m)                               |



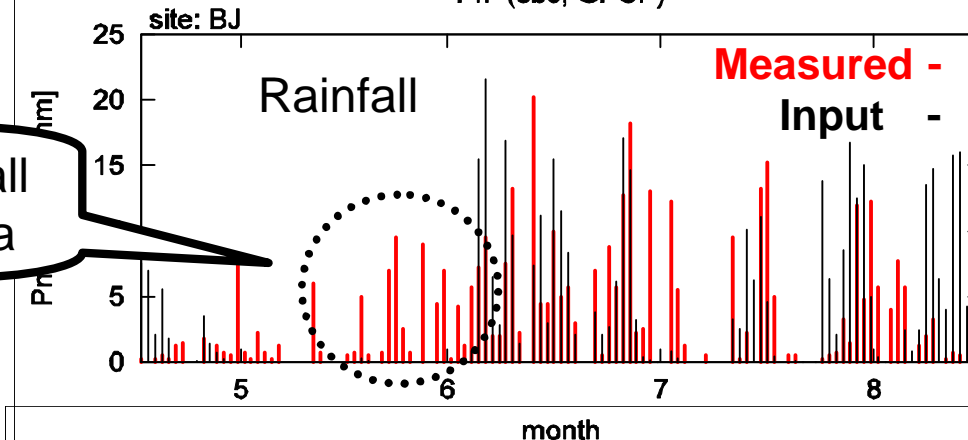
# Comparisons between Assimilation and No assimilation

- Contamination of rainfall errors

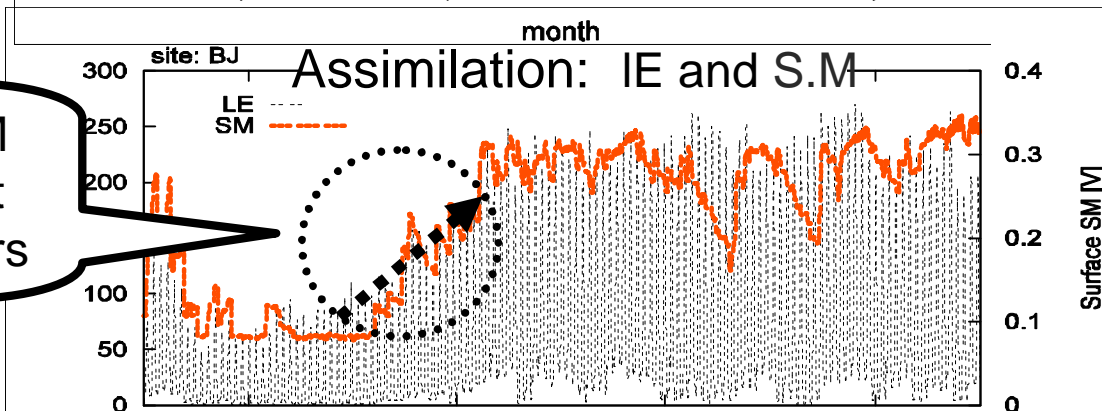
No assimilation: No increase of S.M and IE



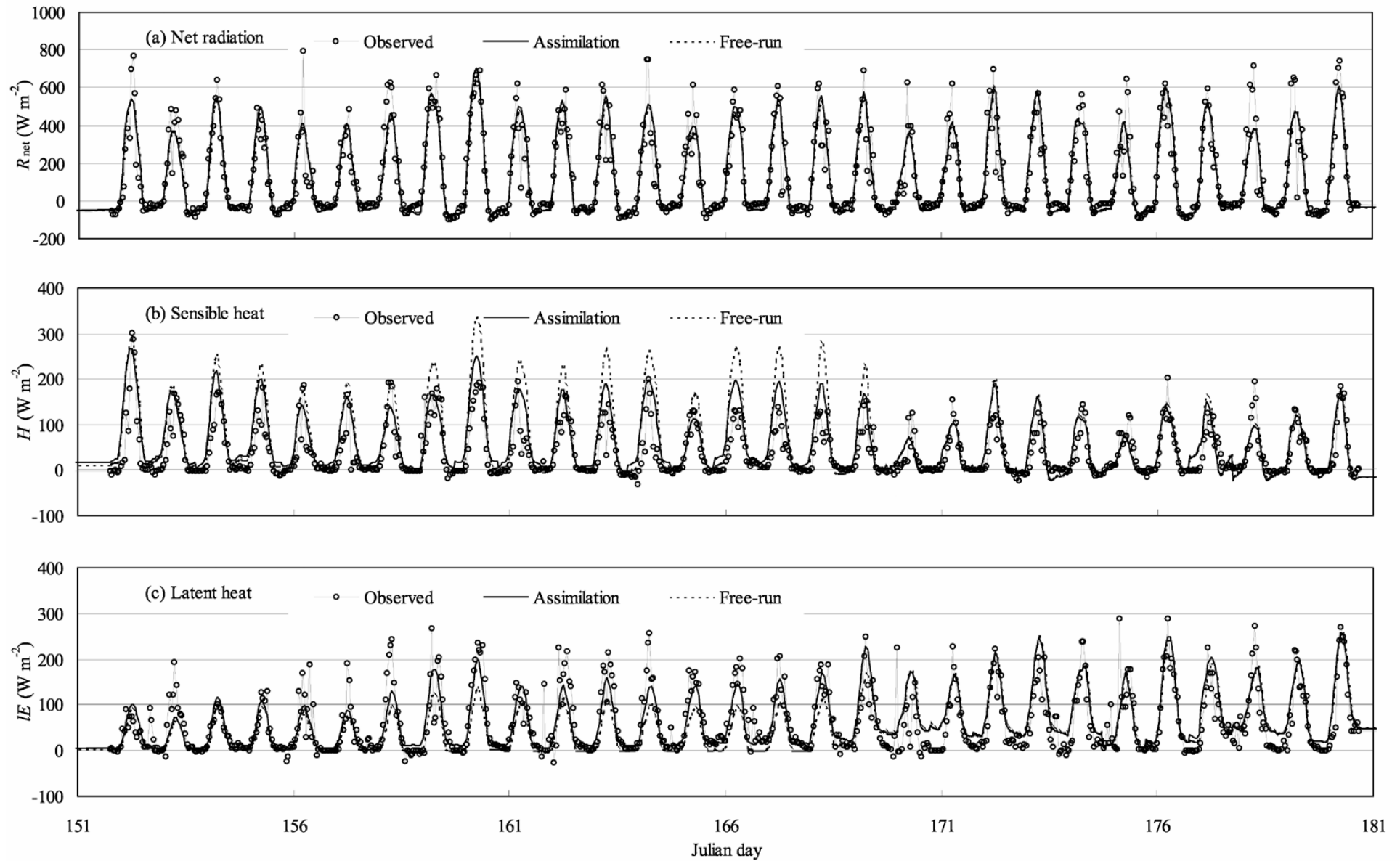
Before June, observed rainfall is not available in GPCP data



Assimilation: Increase of S.M and IE Assimilation is not very sensitive to rainfall errors



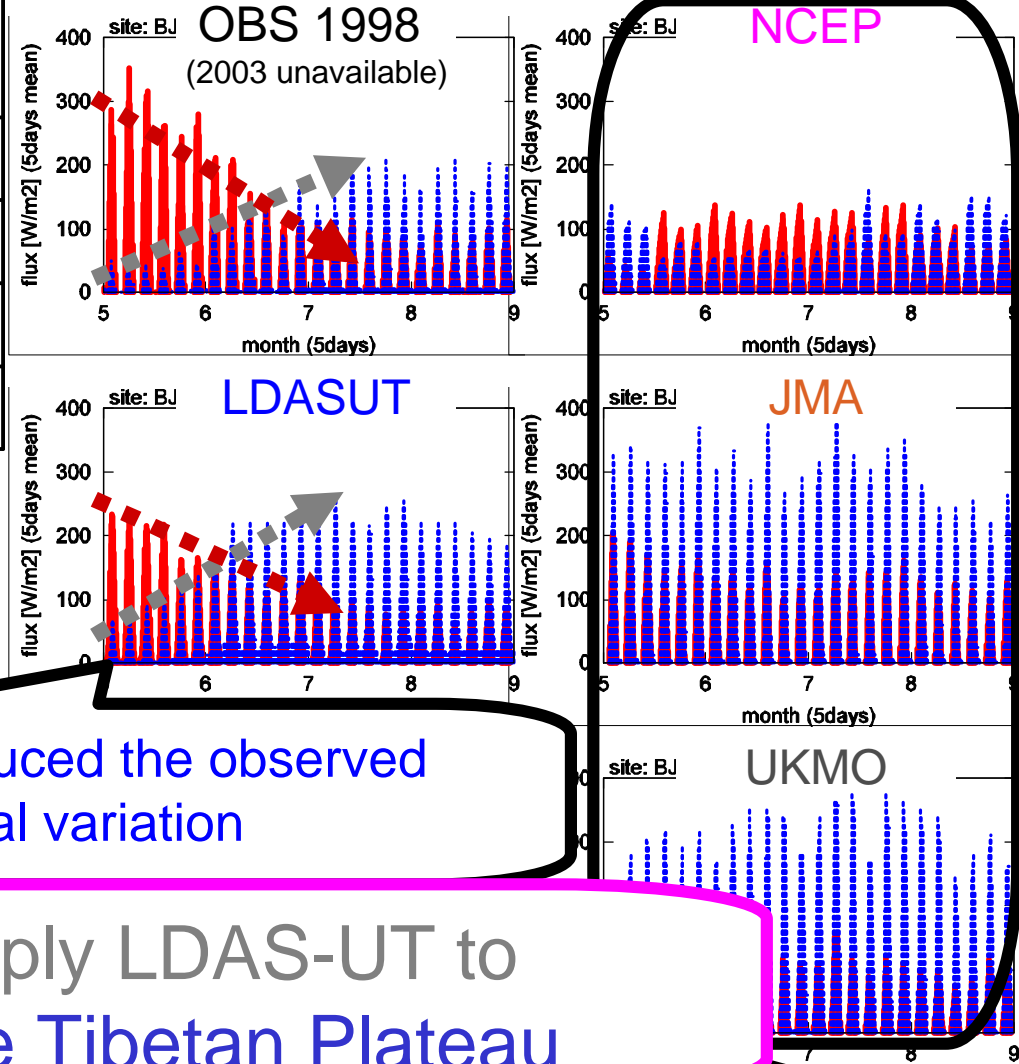
# Surface energy budget



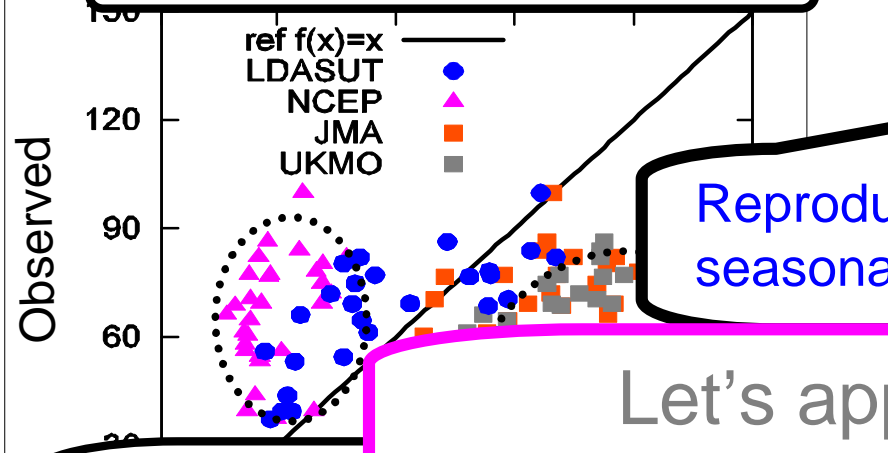
# Compare with GCMs

|               | H RMSE<br>[W/m <sup>2</sup> ] | LE RMSE<br>[W/m <sup>2</sup> ] |
|---------------|-------------------------------|--------------------------------|
| <b>LDASUT</b> | 32.0                          | 42.5                           |
| <b>NCEP</b>   | 40.2                          | <b>68.4</b>                    |
| <b>JMA</b>    | 32.3                          | <b>79.8</b>                    |
| <b>UKMO</b>   | 35.3                          | <b>80.1</b>                    |

Flux variation (May~September)  
**Sensible (H) -**  
**Latent(LE) -**



LE daily-mean ( June )



Reproduced the observed seasonal variation

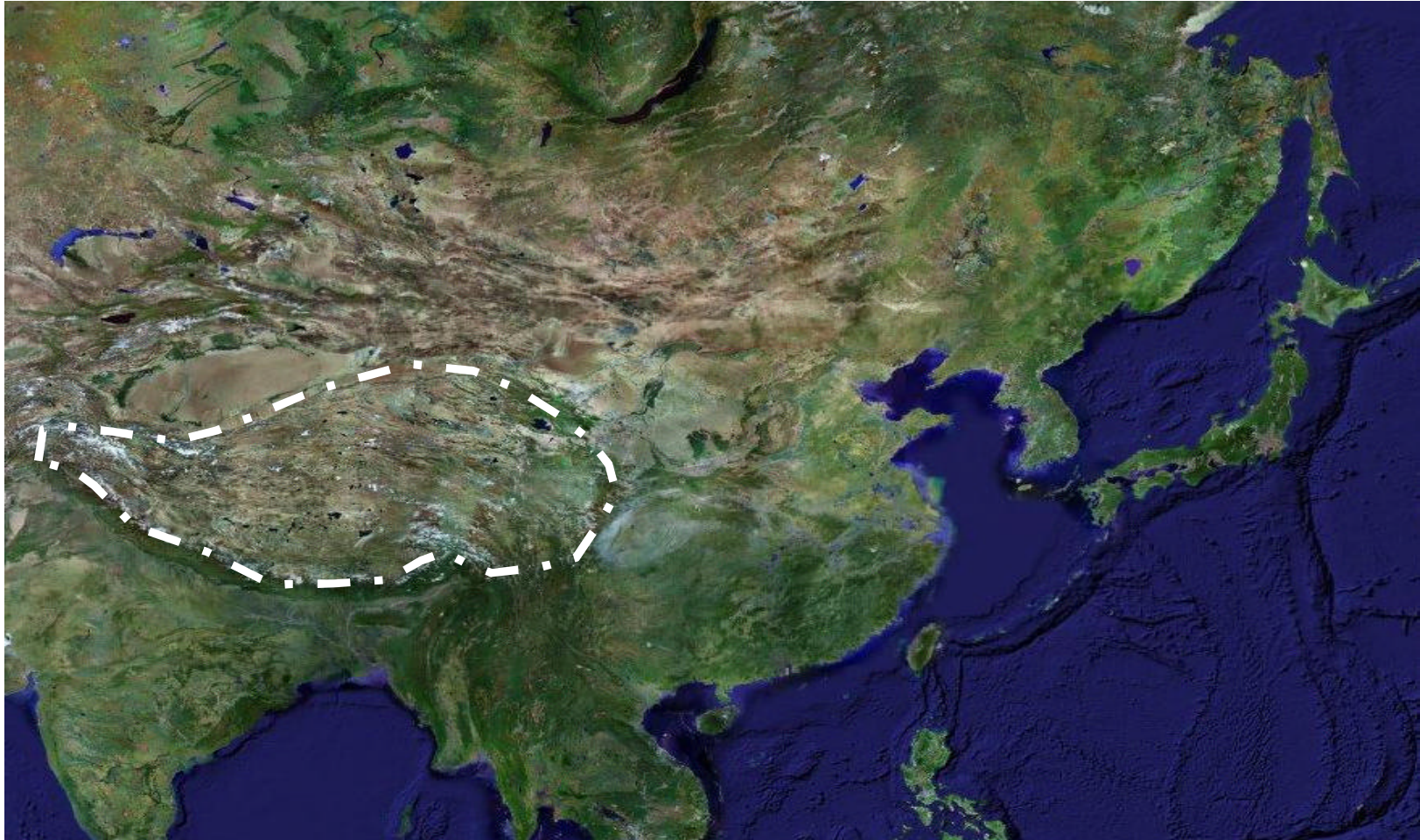
Let's apply LDAS-UT to the whole Tibetan Plateau

LDASUT predicted much better

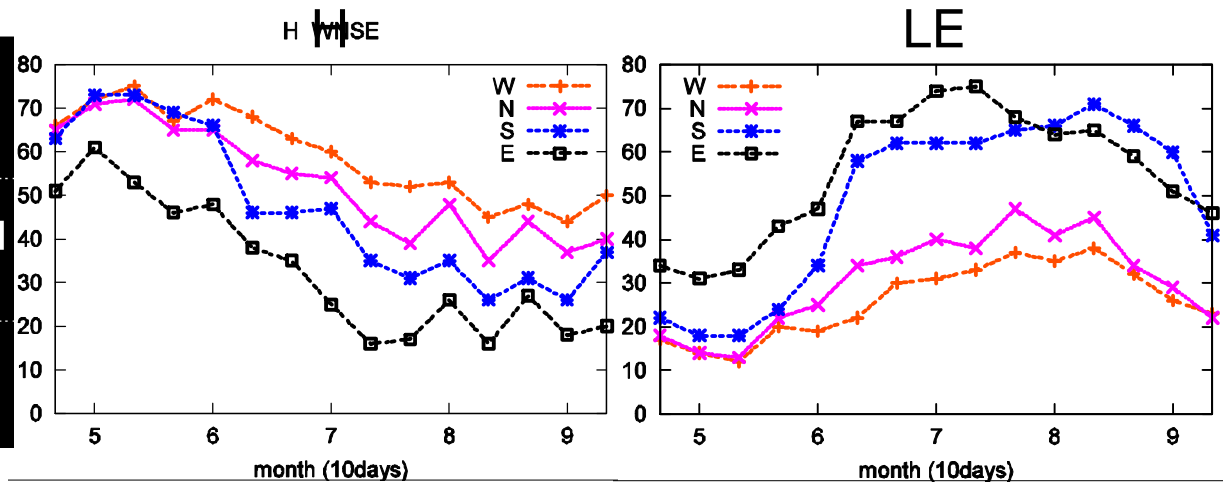
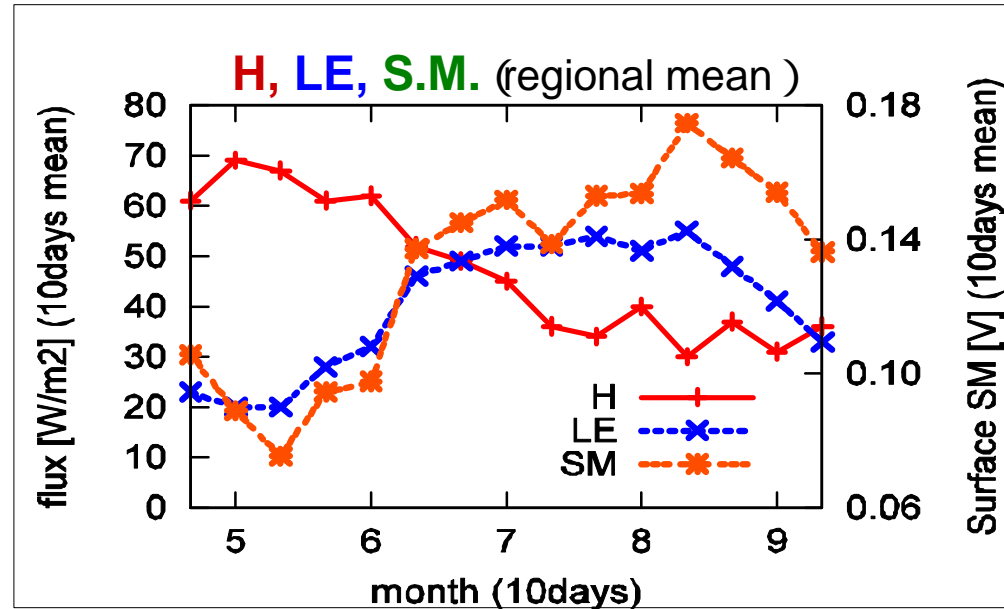
No seasonal variation



# Application to the Tibetan Plateau 2003



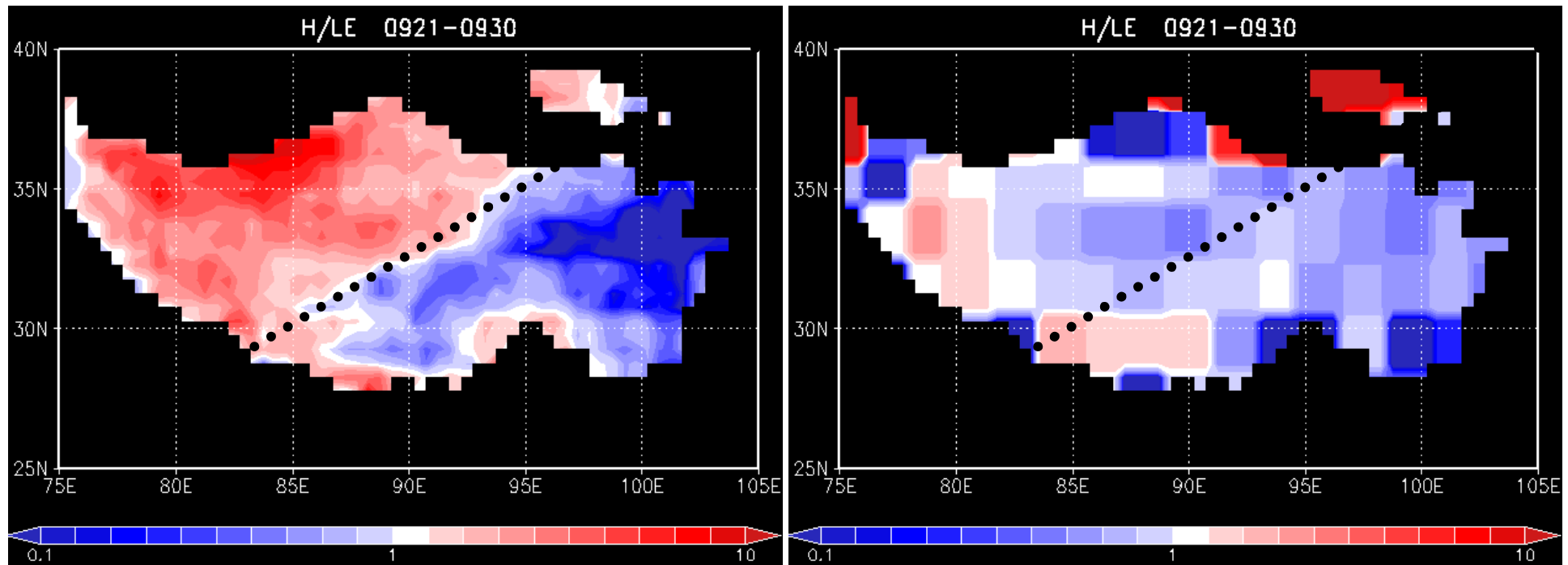
# Seasonality of TP-mean fluxes and S.M.



# Seasonality of distributed Bowen ratio

LDASUT

NCEP



LDAS Seasonality: May~Mid June,  $H > IE$ ; Mid June~Aug;  $IE > H$

LDAS Regionality: H is dominant in N.W. TP, IE is dominant in S.E. TP